

1 **Title: Environmental drivers of sediment carbon storage in temperate seagrass meadows**

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9

10 **Abstract**

11 Seagrass meadows are productive ecosystems that contribute to climate change mitigation by
12 accumulating ‘Blue Carbon’ in their plant biomass and sediments. However, there is wide
13 variation in reported sediment carbon stocks (C_{stocks}) across different global regions and
14 between meadows composed by different seagrass species. Therefore, understanding the
15 drivers for sediment carbon stocks (C_{stocks}) variation is crucial to developing effective
16 conservation and restoration projects for seagrass ecosystems. This study compares and
17 analyses the influence of a range of environmental variables on the variation in sediment C_{stocks}
18 for six intertidal seagrass meadows within the Solent, in southern England. There were
19 significant differences between sites for all variables except leaf density, and concentrations
20 of the sediment pore water nutrients, nitrites and sulphates. Sediment dry bulk density, mean
21 grain size, sorting coefficient, % mud, elevation above sea level, and pore water salinity
22 showed the highest levels of association with C_{stocks} when assessed individually. Multivariate
23 analyses showed that sediment dry bulk density, sorting coefficient, % mud, and pore water
24 pH and concentration of nutrients showed the greatest influence on C_{stocks} . Moreover,
25 sediment characteristics such as dry bulk density, sorting coefficient and % mud, acted in

26 conjunction to explain the bulk of the variation in C_{stocks} among sites. Therefore, sediment
27 characteristics should be considered as important indicators for carbon storage potential in
28 intertidal temperate seagrass meadows used for climate change mitigation.

29

30 **Keywords: Blue carbon, temperate, intertidal, sediment carbon stocks, Seagrass.**

31

32 **Introduction**

33 “Blue Carbon” is a concept that represents carbon stored in coastal ecosystems, including
34 salt marshes, mangroves and seagrass meadows (Pendleton *et al.*, 2012). These ecosystems store
35 carbon within their sediments, living aboveground biomass (leaves, branches, and stems), living
36 belowground biomass (roots), and non-living biomass (e.g., sedimentary organic matter, litter and
37 dead wood) (McLeod *et al.*, 2011). The majority of the blue carbon stored within coastal
38 ecosystems, can be found in their sediments (Murray *et al.*, 2011). These systems are able to
39 sequester and store large amounts of carbon, not only through photosynthesis, but also by trapping
40 sediments and allochthonous organic debris derived from proximate ecosystems and transported
41 by rivers or tides (McLeod *et al.*, 2011; Howard *et al.*, 2017). Whilst mangroves are limited to
42 tropical and sub-tropical zones, salt marshes and seagrass meadows have a pan-global distribution,
43 therefore providing significant carbon storage and sequestration potential (Garrard & Beaumont,
44 2014). It has been suggested that seagrasses provide a higher contribution to carbon accumulation
45 per unit area than terrestrial soils, mainly due to their ability to trap suspended particles by reducing
46 water flow and wave energy (Fonseca & Cahalan, 1992; Gacia & Duarte, 2001; Agawin & Duarte,

47 2002; Gacia *et al.*, 2002; Koch *et al.*, 2006; Bos *et al.*, 2007; Hendriks *et al.*, 2008; Kennedy &
48 Björk, 2009; Rohr *et al.*, 2018).

49 Seagrasses are found in widely varied environments, from sheltered estuaries to highly
50 exposed shores, and from intertidal zones to a depth of 90m in the ocean (Duarte, 1991; Hemminga
51 & Duarte, 2000, Chmura & Hung, 2003; Carruthers *et al.*, 2007; Mazarrasa *et al.*, 2018). This
52 variation in distribution and setting, is one of the factors responsible for the highly variable
53 estimates of sediment organic carbon (C_{org}) and sequestration rates among seagrass communities
54 and species (Lavery *et al.*, 2013; Miyajima *et al.*, 2015; Mazarrasa *et al.*, 2017a). Global
55 estimations report that seagrass sediment carbon storage is believed to average 830 Mg ha⁻¹,
56 resulting in a total estimated global carbon storage of 19.9×10^9 Mg (Fourqurean *et al.*, 2012;
57 Macreadie *et al.*, 2013). However, recent studies have described the influence of habitat
58 characteristics on sequestration and storage of C_{org} in seagrass sediments, which highlights the risks
59 of extrapolating to regional and global estimates from limited data sets represented by only a few
60 species and sites (Nelleman *et al.*, 2009; Fourqurean *et al.*, 2012; Lavery *et al.*, 2013; Garrard &
61 Beaumont, 2014; Serrano *et al.*, 2014; Samper-Villarreal *et al.*, 2016; Serrano *et al.*, 2016; Maxwell
62 *et al.*, 2017; Mazarrasa *et al.*, 2018).

63 The deposition of organic carbon in seagrass meadow sediments (C_{stock}) is regulated by
64 three main mechanisms: meadow productivity and biomass build-up (particularly below-ground);
65 the retention of allochthonous carbon in the sediment; and carbon burial efficiency in seagrass
66 sediments (Mazarrasa *et al.*, 2018). These mechanisms have been reported to be positively related
67 to the anoxic conditions of the sediments, the proportion of clay particles, and of refractory,
68 molecularly complex carbon being stored (Mateo *et al.*, 2006; Serrano *et al.*, 2016; Mazarrasa *et*
69 *al.*, 2018). However, seagrass meadows are experiencing a global area decline estimated at 7% per

70 year, potentially resulting in CO₂ emissions as sediments are increasingly being eroded and C_{stocks}
71 exposed to aerobic conditions (Waycott et al., 2009; Marbà *et al.*, 2014; Serrano *et al.*, 2016;
72 Lovelock *et al.*, 2017; Mazarrasa *et al.*, 2018). Santos *et al.*, (2019) highlighted similar trends in
73 seagrass decline in Europe, despite efforts to include these ecosystems as sensitive quality elements
74 to provide diagnostic of ecosystem health under the European Union (EU) Water Framework
75 Directive (WFD). The decline in seagrass C_{stocks} is due to a range of anthropogenic impacts related
76 to eutrophication, shading, shoreline erosion, sea warming, and physical removal of shoots by
77 trawling and anchoring (Duarte, 2002; Orth *et al.*, 2006; Ralph *et al.*, 2006; Macreadie *et al.*, 2012;
78 Marbà *et al.*, 2013; Duarte, 2014; Santos *et al.*, 2019). Conversely, studies have shown that
79 eutrophication and nutrient loading might favor the accumulation of allochthonous carbon (e.g.
80 microalgae and epiphyte blooms) in sediment deposits, leading to an increase in the total carbon
81 sequestered in seagrass meadows (Macreadie *et al.*, 2012; Serrano *et al.*, 2016; Mazarrasa *et*
82 *al.*, 2017b; Samper-Villarreal *et al.*, 2018).

83 Furthermore, climate change may aggravate seagrass decline, as a result of multiple impacts
84 including ocean acidification, and increases in sea surface temperature and water depths (Short &
85 Neckles, 1999; Jordà *et al.*, 2012; Saunders *et al.*, 2013; Valle *et al.*, 2014; Marba *et al.*, 2018). The
86 effects of ocean acidification on the accumulation of autochthonous C_{org} in seagrass sediments and
87 biomass still needs to be clarified, with some studies suggesting increases in sequestration rates
88 under acidic conditions (Palacios & Zimmerman, 2007; Hall-Spencer *et al.*, 2008; Fabricius *et al.*,
89 2011; Russell *et al.*, 2013; Garrard & Beaumont, 2014; Mazarrasa *et al.*, 2018) and others reporting
90 a substantial decrease (Martínez-Crego *et al.*, 2014; Repolho *et al.*, 2017). Thus, understanding the
91 response of seagrass ecosystems to climate change and other anthropogenic impacts has become a
92 priority for the development of effective conservation and management (Brierley & Kingsford,
93 2009; Hoegh-Guldberg & Bruno, 2010; Valle *et al.*, 2014; Maxwell *et al.*, 2017; Santos *et al.*,

94 2019). It is important to draw attention to the potential that environmental factors can act in
95 synergy, and their effects will depend upon seagrass species' abilities to tolerate and adapt to
96 different scenarios (Ralph, 1999; Sunda & Cai; 2012; Repolho *et al.*, 2017). The high sensitivity
97 of seagrasses to environmental change means it is vital to understand which factors may threaten
98 their role as long term carbon sinks (Marbà *et al.*, 2012; Cullen-Unsworth *et al.*, 2014; Jones &
99 Unsworth, 2016; Mazarrasa *et al.*, 2018). Therefore, monitoring programs based either directly or
100 indirectly on seagrass responses to environmental disturbance have increasingly been incorporated
101 into ecosystem management (Martínez-Crego *et al.*, 2008; Montefalcone, 2009; Roca *et al.*, 2016).
102 For example, the European Union's Water Framework Directive includes the monitoring of
103 seagrass ecosystems as a reference for the ecological status of coastal areas, using characteristics
104 such as leaf density, cover, and depth limits (Longstaff *et al.*, 1999; D'Souza *et al.*, 2015; Roca *et*
105 *al.*, 2016).

106 A large number of biotic and abiotic factors can potentially influence carbon stocks and
107 accumulation rates in seagrass meadows (Maxwell *et al.*, 2017). Mazarrasa *et al.*'s (2018) review
108 found that species composition, high canopy complexity, a continuous meadow landscape, biotic
109 interactions made of complex and stable trophic interactions, low exposure to wave energy, low
110 levels of turbidity, and shallow water depth, were key factors contributing favourably to carbon
111 storage. Conversely, factors that negatively impact carbon storage were, low nutrient availability,
112 over grazing, bioturbation, eutrophication and climate change, while elevation, climatic region and
113 acidification were amongst the unresolved components (Maxwell *et al.*, 2017; Mazarrasa *et al.*,
114 2018). However, most studies investigating relationships between environmental factors and
115 seagrass carbon storage capacity have focussed on specific factors individually, rather than their
116 associations (Lavery *et al.*, 2013; Rozaimi *et al.*, 2013; Duarte *et al.*, 2013; Martínez-Crego *et al.*,
117 2014; Armitage and Fourqurean 2016; Howard *et al.*, 2016; Ricart *et al.*, 2017; Oreska *et al.*, 2017;

118 Mazarrasa *et al.*, 2017; Mazarrasa *et al.*, 2018). Few studies have analysed the influence of multiple
119 environmental factors as drivers of variability in carbon storage in seagrass sediment, with some
120 evaluating relationships at small (within meadow) scales (Samper-Villarreal *et al* 2016; Mazarrasa
121 *et al.*, 2017), and others comparing between meadows from different geographical regions (Lavery
122 *et al.*, 2013; Miyajima *et al.*, 2015; Dahl *et al.*, 2016; Gullström *et al* 2017; Rohr *et al.*, 2018),
123 which could make a reliable assessment of variability harder.

124 In addition, intertidal seagrass meadows may be particularly vulnerable to multiple stressors, such
125 as air exposure, temperature, light intensity and salinity, which could impact photosynthetic rates
126 and consequent carbon uptake and storage (Bjork *et al.*, 1999). Intertidal populations are also prone
127 to runoff from catchment areas, being susceptible to elevated levels of nutrients from industrial and
128 agricultural waste, which not only affect meadow health, but increase epiphyte productivity (Short
129 and Willie-Echeverria, 1996; Ye *et al.*, 2003). Variability among seagrass species can also affect
130 their carbon storage potential, with species-environment interactions likely to strongly impact the
131 storage of sedimentary carbon (Lavery *et al.*, 2013). To this date, no study has evaluated the
132 interaction between environmental factors and variation in seagrass sediment carbon storage at the
133 intermediate (regional) scale, incorporating different habitat characteristics, and mixed species
134 meadows. Further understanding of these interactions would benefit global seagrass research,
135 promoting conservation and protection programs with the aim of climate change mitigation, and
136 seagrass status as a biological indicator of ecosystem status. Therefore, this study aims to identify
137 the key environmental factors driving carbon storage in intertidal seagrass sediments from different
138 sites within the same temperate region. In this study, i) variation in environmental factors between
139 seagrass sites is assessed, and ii) relationships between factors that influence carbon stocks were
140 identified, to answer the following research hypothesis: within temperate intertidal seagrass
141 meadows, sediment characteristics most strongly influence carbon storage.

142

143 **Methods**

144 **Study Sites**

145 The Solent is considered one of the most important coastal areas in the UK composed of
146 natural and man-made environments with high habitat diversity, providing an important wildlife
147 resource internationally (King, 2010). The distinctive hydrographic regime of an extended or
148 double high tide and the intricacy of different habitats are reasons why the Solent region has been
149 selected as a Special Area of Conservation (SAC) (McLeod *et al.*, 2005). Tidal amplitudes are not
150 uniform along the Solent, with the eastern end having almost double the tidal range of the western
151 end, providing longer inundation periods (Dyer & King, 1975). The region is defined by the body
152 of water that encompasses the central south coast of England and the Isle of Wight, thought to have
153 been formed during the last Ice Age, when ice melting flooded the paleo-channel Solent River
154 system (Fletcher *et al.*, 2007).

155 Studies have been reporting water quality issues along the Solent (Harding *et al.*, 2016;
156 Environment Agency 2016a; 2016b). For example, the harbours and estuaries at Langstone,
157 Chichester, Portsmouth, and Eastern Yar, on the Isle Wight, have been classified as eutrophic or at
158 risk of eutrophication by the Environment Agency (Harding *et al.*, 2016; Environment Agency
159 2016a; 2016b). This increase in nutrient levels can cause algal blooms, promoting the growth of
160 benthic algae which can potentially smother seabed habitats, like seagrass meadows (Harding *et*
161 *al.*, 2016). Therefore, these estuaries in the Solent have been designated as sensitive areas, or
162 polluted waters, under the Urban Waste Water Treatment Directive and/or Nitrates Directive
163 (Harding *et al.*, 2016; SeaView, 2017). Sources of pollution include treated effluent discharges
164 from waste water treatment works, and runoff from the surrounding catchment including waste and

165 industrial activities (SeaView, 2017). Combined with riverine sources, it is evident that offshore
166 coastal background sources have been responsible for the nutrient loading input within the Solent's
167 estuaries (Environment Agency 2016a; 2016b).

168 To represent the characteristic variation in the region, six study sites with known seagrass
169 meadow extents, were selected within the Solent region, Southeast of England, following an
170 assessment of the most recent seagrass distribution inventory (Marsden and Chesworth, 2015)
171 (**Table 2**). Selected sites were Creek Rythe in Chichester Harbour, Hayling Island and Farlington
172 Marshes in Langstone Harbour, Porchester in Portsmouth Harbour, Cowes and Ryde (**Figure 1**).
173 Seagrass meadows from all sites are characterised as intertidal, located in both sheltered inland
174 bays (Chichester Harbour, Langstone Harbour, and Portsmouth Harbour sites), and more exposed
175 shorelines (Isle of Wight sites), encompassing seagrass habitats from both muddy and sandy
176 substrates, incorporating *Zostera marina* (Eelgrass), *Zostera noltei*, *Zostera angustifolia*, and
177 *Ruppia spp.* meadows (Marsden & Chesworth, 2015). The most recent seagrass surveys conducted
178 at Creek Rythe, in Chichester Harbour, reported patchiness with varied meadow density, and no
179 clear dominance between *Z. noltei* and *Z. angustifolia* species, while *Ruppia spp.* were only found
180 within inlet channels (Marsden & Chesworth, 2015). The site has a tidal range of 0.9-4.9m,
181 represented by mean low water springs (MLWS) and mean high water springs (MHWS) values,
182 respectively. Surveys in Langstone Harbour and Porchester reported mainly the presence of *Z.*
183 *angustifolia* and *Z. noltei*, but also *Ruppia spp.* in intertidal areas (Marsden & Chesworth, 2015),
184 with a MLWS and MHWS tidal range of 0.8-4.8m, respectively. Here, there are reports of
185 significant declines in seagrass due to extensive trampling and dredging, and some evidence of
186 anoxic conditions and smothering from dense green algae mats (Marsden & Chesworth, 2015). At
187 the Isle of Wight sites, Cowes is characterised by gravel and soft to firm sandy sediments, with

188 intertidal-subtidal seagrass coverage to 2m below Chart Datum and MLWS and MHWS tidal range
189 between 0.8-4.2m, while Ryde is characterised by soft to firm sandy sediments (Marsden &
190 Chesworth, 2015), with a MLWS and MHWS tidal range of 0.2-3.1m, respectively. Surveys at both
191 sites reported the presence of *Zostera* spp., including *Z. marina* and *Z. noltei*, with possible *Z.*
192 *angustifolia* at Ryde (Marsden & Chesworth, 2015).

193 **Field Methods**

194 A range of environmental variables were measured at each study site to determine factors
195 influencing C_{stocks} (**Table 2**). Sediment variables included grain size, dry bulk density, degree of
196 sorting and % mud. Pore water parameters were salinity, pH, and nutrients (nitrites $[\text{NO}_2^-]$ and
197 sulphates $[\text{SO}_4^{2-}]$). Elevation above sea level and the biological parameters leaf density and above-
198 ground biomass were also recorded.

199 Degree of sorting, calculated as the sorting coefficient from the different sediment grain size

Fig. 1 Location of the six seagrass sampling sites in the Solent, southern England (red square). Zoomed image shows seagrass sampling sites and their respective seagrass meadows areal extent in red, collated by Marsden and Chesworth (2015): Hayling Island (100.24 ha), Creek Rythe (70.1 ha), Porchester (94.92 ha), Farlington Marshes (31.2 ha), Cowes (27.1 ha) and Ryde (82.47). Maps are adapted from Esri ArcGIS online basemaps

200 fractions, was used along elevation above mean sea-level as a proxy for degree of exposure of the
201 site (see Folk & Ward, 1957). Mud content (% mud), including clay and silt particles ($< 63 \mu\text{m}$),
202 has been suggested as a better representative fraction of seagrass bulk sediment and their C_{stocks} ,
203 than solely using clay particles ($< 4 \mu\text{m}$) (De Falco *et al.*, 2004; Burdige, 2007; Pedrosa-Pamies *et*
204 *al.*, 2013; Serrano *et al.*, 2016). Therefore, %mud ($> 63 \mu\text{m}$) was selected for analysis, including
205 both silt and clay fractions.

206 Field sampling was conducted during low tide, when seagrass meadows were exposed. Five
207 sampling points were selected within each of the six study sites, thirty sampling points in total
208 (Howard *et al.*, 2014). Sampling points were randomly selected by walking towards the middle of

209 the meadow, at least 3m from the edge, and randomly throwing 0.25m² quadrats in a clockwise
210 orientation. From each of the thirty randomly selected sampling points, the following samples were
211 collected during the summer (June-August) of 2017: above-ground biomass and leaf density in the
212 form of cropped leaves within a 0.25m² quadrat; one, 30cm deep, sediment core for mean pooled
213 carbon stocks and particle size analyses; one, 30cm deep, sediment core for pore water analyses.
214 A dGPS (Leica GPS1200 Surveying System) was used to record latitude, longitude, and elevation
215 above mean sea level (vertical accuracy, 0.02 m) from ten points within each sample quadrat (50
216 points per site; 300 points in total)(Ward *et al.*, 2016). Data were post processed using the Leica
217 Geo Office software version 8.4, correction data available from the RINEX (Receiver Independent
218 Exchange Format), downloaded from Leica Geosystems using the British National Grid coordinate
219 system. This was plotted using the OSGB36 datum, and the reference station used was Sandown,
220 Isle of Wight (50°39'5.69" N -1°09'39.71" W).

221 Seagrass species were identified, and a visual assessment of meadow landscape, very
222 patchy (< 20% cover), patchy (20 ≤70% cover) or dense (> 70% cover), was conducted with a
223 walkover of the sites. Above-ground biomass and leaf density was recorded from each sample plot
224 by cropping the plant biomass (leaves to stem base) within each 0.25 m² quadrat, before being
225 stored at -20 °C prior to analysis (Howard *et al.*, 2014). Cores were collected using a Russian corer,
226 with a 5cm diameter, to avoid core compaction. Each core was divided into 5cm depth subsamples,
227 with 6 subsamples per core (Howard, *et al.*, 2014). Because the oxygenate rhizosphere layer is
228 more likely to be affected by changes in environmental conditions, due to microbial activity and
229 sediment deposition processes (Enriquez *et al.*, 2001; Gray and Elliot, 2009), cores were taken in
230 the top 30cm layer of sediment, or until refusal (20cm at Cowes). Post collection, the 30 sediment
231 cores collected for carbon stocks analysis were kept in a freezer at -20°C, and the remaining 30

232 sediment cores, used for pore water extraction, were kept in a cold storage room at 4°C at the
233 University of Brighton’s sediment analysis lab, for < 48 hours.

234 **Laboratory methods**

235 In the laboratory, above-ground biomass was transferred to 1mm sieves, and washed free of
236 sediment under running water to separate living above-ground components (Howard, *et al.*, 2014).

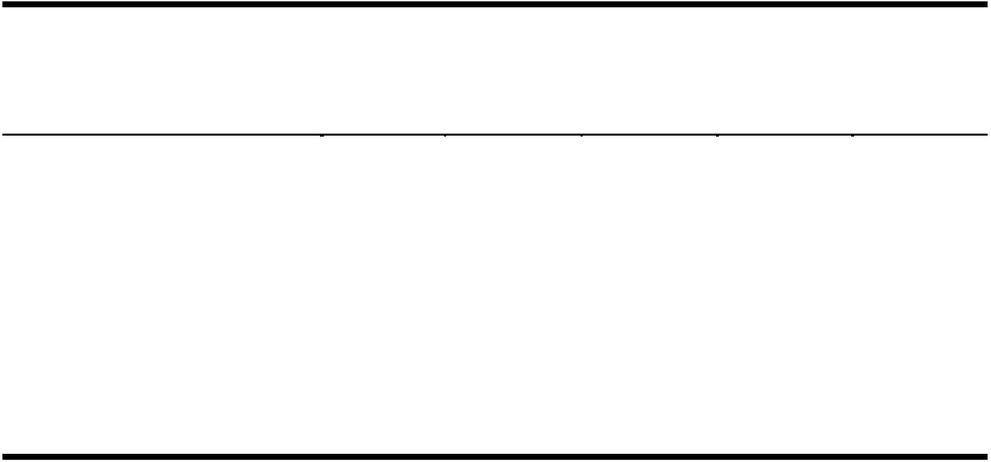
237 The identified seagrass species were recorded and whole leaves (stem to tip) were counted from
238 each sample to determine leaf density. Filamentous macroalgae and invertebrates were separated
239 from seagrass biomass during the washing procedure, however, for the sake of this study,
240 microalgae epiphytic load, when found, was not scraped from the leaves, to prevent loss of
241 vegetative organic matter. Above-ground plant biomass was determined by oven-drying the
242 vegetative biomass to a constant weight (72 h at 60 °C) (Howard, *et al.*, 2014). The above-ground
243 living vegetative component was determined by multiplying the dry weight (kg) of a sample of
244 plant material for a given area (m²) by a carbon conversion factor (0.34), derived from literature
245 for seagrass above-ground biomass calculations (Duarte, 1990; Howard *et al.*, 2014).

246 After thawing, each sediment subsample was weighed prior to oven drying at 60 °C for at least 72
247 hours, and then cooled at room temperature in a desiccator for at least one hour before weighing
248 again to determine moisture content (Howard *et al.*, 2014). Oven dried sediment samples were
249 disaggregated with a pestle and mortar and weighed in individual beakers, 2-4g for each sample,
250 prior to analysis of organic matter. To estimate sediment organic carbon, sequential Loss on
251 Ignition (% LOI) at 450 °C for 24h was selected as this method has been found to correlate well
252 with estimation of C_{org} in seagrass meadows (Fourqurean *et al.*, 2012b; Macreadie *et al.*, 2014).

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Samples were cooled at room temperature in a desiccator for at least one hour before weighing to determine percentage of organic matter (% OM) (Heiri *et al.*, 2001). To determine sediment C_{stocks}, dry bulk density, organic carbon content (C_{org}), and carbon density were calculated for each subsample. Dry bulk density (g/cm³) for each subsample was estimated using the equation by Dadey *et al.*, 1992:

$$\text{Equation 1: } Pd = (1 - \phi) * Ps$$

Where Pd = bulk density, ϕ = porosity, and Ps = grain specific gravity.

%C_{org} was determined using regression equations derived from the literature for seagrass (Fourqurean *et al.*, 2012a/b):

$$\begin{aligned} \% C_{org} &= 0.43 * \% OM - 0.33, \text{presuming } \% OM > 0.2 \\ \% C_{org} &= 0.40 * \% OM - 0.21, \text{presuming } \% OM < 0.2 \end{aligned}$$

279
280 Following %C_{org} calculations, carbon density and carbon content were determined using the
281 equations described by Howard *et al.* (2014). C_{stock} results from each subsample were summed to
282 determine total carbon in each core and converted to Mg C/ha.

283 Following LOI, particle size analysis was carried out on all non-ground sediment samples using a
284 Malvern Mastersizer 2000 laser analyser, with particle size grading undertaken in accordance with
285 the Wentworth (1922) size classification scheme. Samples were washed with 10 ml of sodium
286 hexametaphosphate prior to analysis and then stirred for 5 minutes in order to deflocculate clay
287 particles (Ward *et al.*, 2014). A small sub-sample ~1 – 1.5 g, dependent on laser obscuration related
288 to particle size, was analysed using a basic ultrasonic setting, which improves dispersion of
289 particles during analysis by breaking up aggregates using vibrating sound waves (Malvern
290 Instruments, 1980). The final data for each size classification (clay, silt, and sand), represented an
291 average of three separate analytical runs (standard error < 1 %) (Ward *et al.*, 2014).
292 The mean (central value), and sorting coefficient (standard deviation) were calculated for each
293 sample following Folk & Ward's method (1957).

294 Each 5cm increment from the 30 remaining sediment cores kept in cold storage were
295 divided into two subsamples, one for analysis of the concentration of nitrites and sulphates, and
296 one for pH and salinity. Sediment pore water has been described as the main provider of nutrients
297 for seagrass growth, being several orders of magnitude higher in concentration than nutrients in the
298 water column (Fourqurean *et al.*, 1992; McGlathery *et al.*, 2001). Thus, pore water was extracted
299 from each sediment subsample within 48h of collection, to prevent organic decomposition
300 (Michalski and Kurzyca, 2005; EPA, 2007). For pore-water extraction, each sediment subsample
301 was centrifuged using an Eppendorf™ 5702 Series Centrifuge for 15 min at 4,400 rpm.
302 Supernatant was collected to perform dilution trials and determine the most suitable dilution factor

303 to better identify relevant peaks. Dilution ratios of 1:100; 1:10; 1:5 and 1:2 were tested, between
304 the extracted pore water and deionised water, adding up to a total volume of 5ml (Jackson, 2000).
305 Due to the instability of the nitrogen oxide ions, and problems related to pairing separations of Cl^-
306 $/\text{NO}_2^-$ in saline samples (Michalski and Kurzyca, 2005), only nitrites and sulphates peaks were
307 clearly detected (using a dilution factor of 2), and the concentration of both anions was converted
308 from mg/L into μM prior to analysis.

309 Salinity and pH of each pore water subsample was measured, following the 2:5 ratio
310 proposed by Head (2006). Approximately 3g of dry sediment sample was mixed with 7.5ml of
311 distilled water in a temperature controlled orbital shaker for 10 min to dissolve particles for analysis
312 (Head, 2006). Samples were allowed to stand overnight and stirred again immediately before
313 testing (Head, 2006). pH analysis was conducted using a Mettler Toledo™ FE20 FiveEasy™
314 Benchtop pH Meter. Tests were conducted with three replicates, stirring briefly between readings
315 to ensure accuracy, and an average was calculated. Probes were washed with distilled water
316 between tests and dried before use (Head, 2006; Burnside *et al.*, 2008). Equipment was calibrated
317 using buffer solutions of pH 4.0 and 7.0. Supernatants from the same (previously stirred) samples
318 used for pH analysis were used to determine salinity. Small droplets were applied to a Bellingham
319 + Stanley™ Eclipse Hand Held Refractometer 45-63, to measure salinity (‰). Three replicates of
320 each reading were measured and an average was calculated. The refractometer was calibrated
321 between cores, by taking a reading using distilled water.

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324 **Statistical Analysis**

325 All statistical analysis was performed on Minitab 2017. Normality probability plots and
326 histogram frequency of residuals were tested using the theory-driven Anderson-Darling method for
327 each of the 12 variables (Anderson and Darling, 1954): pooled sediment C_{stock} (30cm), dry bulk
328 density [DBD], mean grain size, sorting coefficient, %mud, pH, salinity, nitrite (NO_2^-), sulphate
329 (SO_4^{2-}), elevation, leaf density and above-ground biomass. When assumptions of normality were
330 not met, variables were transformed using $\text{Log}_{10}(X)$ and retested for normality (**Supplementary**
331 **Table A**). Site differences between each variable were tested using ANOVA Post-hoc Tukey's test
332 was used to identify significantly different means for each variable, when present. Means are
333 expressed as mean values for each variable from the sampled quadrats, 30 (5 per site) samples in
334 total, to assess for the relationships between each variable and pooled carbon stocks (30cm deep).
335 Individual relationships between variables and sediment C_{stocks} were examined using the following
336 tests. A linear regression was used to derive an equation to determine sediment C_{stocks} based on
337 DBD values as predictors, since DBD is commonly used in standard calculations of sediment
338 carbon density (Howard *et al.*, 2014). A Pearson's correlation test, or where data were not normally
339 distributed a Spearman's Rho correlation, was used to assess the relationship between variables
340 and C_{stocks} . A Partial least squares (PLS) regression multivariate analysis was used to assess the
341 influence of different types of factors on sediment C_{stocks} (Dahl *et al.*, 2016; Gullström *et al.*, 2018;
342 Rohr *et al.*, 2018). The advantage of using PLS as a model for multivariate regression is that it can
343 tolerate collinear explanatory factors, with a large number of predictors that might not be fixed, or
344 that might contain measurement errors (Carrascal *et al.*, 2009; Dahl *et al.*, 2016). Results from the
345 PLS analysis were used to identify variables with higher correlation coefficients, showing stronger
346 relationships with sediment C_{stocks} . These variables were analysed by principal component analysis
347 (PCA) to better visualise and understand their level of association, by aggregating and summarising

348 groups of highly inter-correlated variables and explaining the variation in C_{stocks} among sites (ter
349 Braak, 1995; Marin-Guirao *et al.*, 2005). All results assume a significance of $p < 0.05$.

350

351 **Results**

352 **Table 2:** Environmental data for each seagrass study site in the Solent region, UK. Sites are presented in
353 decreasing order of mean sediment C_{stock} . Values are presented as mean (\pm SD), $n=30$ for all variables, with
354 Tukey's grouping results following ANOVA where the same letters correspond to means that are not
355 statistically different, for each variable.

356

357 **Variation among sites**

358

359 The mean C_{stock} within the seagrass sediment across all six sites was 33.80 ± 18.40 MgCha
360 ¹ ($n=30$). However, there was a significant variance in mean sediment carbon storage (30cm)
361 between sampling sites, with those on the Isle of Wight, Cowes and Ryde, having significantly
362 lower sediment C_{stock} values than all other sites, but not significantly different between themselves
363 (**Table 2**). The mean DBD values for all sites was 0.99 ± 0.03 g/cm³ ($n=30$). The mean sediment
364 DBD values at Ryde and Cowes were significantly higher than all other sites (**Table 2**). All sites
365 were classified as muddy, silt rich, sediments according to the Wentworth scale, apart from Cowes
366 and Ryde, which contained average grain size (μ m) within the class of very fine and fine sand
367 respectively. The sorting coefficient at Ryde was significantly higher than all other sites (**Table2**).
368 Porchester and Creek Rythe presented the lowest sorting coefficients respectively, both
369 significantly different than Cowes. However, all sites were classified as very well sorted. Ryde had
370 by far the lowest mean %mud compared to all other sites (1.14%). Hayling Island had the highest
371 mean %mud (93.85%), followed by Creek Rythe (87.50%), both significantly higher than Cowes
372 and Ryde (**Table 2**).

373 The average pore water pH across all sites was 7.28 ± 0.28 (n=30) with relatively little, but
374 significant, variation across all sites (**Table 2**). Mean salinity showed significantly lower levels in
375 the Isle of Wight's sites when compared against all others. Ryde had the lowest sediment pore
376 water salinity (1.4 ± 0.55 ‰), followed by Cowes (2.0 ± 0.0 ‰), whilst the highest values were
377 found in Hayling Island (15.6 ± 0.89 ‰) (**Table 2**). Mean pore-water concentrations of nitrites
378 and sulphates showed no significant differences between sampling sites (**Table 2**). However, there
379 was large variation in concentration of both nutrients with depth within cores, with both sites in the
380 Isle of Wight only presenting detectable levels of pore water nutrients in the upper layers, Cowes
381 to 10cm and Ryde to 15cm depth (**Figure 2**). Cowes showed the largest decrease in concentration
382 of both pore water nutrients, however, Ryde presented an increase in nitrite concentration and a
383 decrease in sulphate levels with depth (**Figure 2**). Down core variation in the concentration of
384 both nutrients was the same in Farlington Marshes and Porchester, but did not follow the same
385 pattern in Creek Rythe and Hayling Island (**Figure 2**).

Fig. 2 Concentrations of (a) nitrites NO_2^- ($\mu\text{mol/L}$) and (b) sulphates SO_4^{2-} ($\mu\text{mol/L}$) down-core for all study sites: Cowes, Ryde, Creek Rythe, Hayling Island, Porchester and Farlington Marshes

386 Seagrass meadows at Cowes were located at a significantly lower elevation above mean
387 sea level than all other sites, at $-1.82 \pm 0.05\text{m}$. (**Table 2**). Out of all sites, only Farlington Marshes
388 and Creek Rythe presented positive values of elevation above mean sea level, significantly different
389 than all other sites, of $0.06 \pm 0.07\text{m}$ and $0.0008 \pm 0.06\text{m}$, respectively. The elevation of the three
390 remaining study sites (Hayling Island, Porchester and Ryde) were all significantly different to each
391 other (**Table 2**).

392 Average leaf density across all sites was 394 ± 268 leaves/ m^2 (n=30). There was no
393 significant difference in leaf density between sampling sites, but large standard deviations were
394 recorded (**Table 2**). *Z. angustifolia* was the dominant species in Farlington Marshes and Ryde, and

395 was present at all sites apart from Cowes. At Farlington Marshes and Cowes, seagrass meadows
396 presented very patchy landscape patterns, with un-vegetated patches within the seagrass beds. At
397 Ryde and Porchester, the landscape was characterised as patchy, with less un-vegetated space
398 between seagrass beds, while Hayling Island and Creek Rythe supported dense beds with
399 continuous seagrass meadows (**Table 2**). There were significant differences in mean above-ground
400 biomass between sampling sites, with Creek Rythe presenting significantly higher above-ground
401 biomass values, of $0.497 \pm 0.25 \text{ MgCha}^{-1}$ (n=5), than both sites on the Isle of Wight (**Table 2**).
402 Ryde had the lowest above-ground biomass values amongst all sites, of $0.07 \pm 0.03 \text{ MgCha}^{-1}$ (n=5),
403 significantly lower than Creek Rythe and Hayling Island (**Table 2**).

404

405

406 **Relationships between environmental variables and sediment C_{stock}**

407 A regression analysis demonstrated a significant negative relationship between C_{stock} and
408 DBD ($R^2 = 87.7\%$, $p = 0.000$) (**Figure 3a**). There was a significant negative relationship between
409 mean grain size and sediment C_{stock} ($r = -0.712$ and $p = 0.000$) (**Figure 3b**), and between sorting
410 coefficient and sediment C_{stocks} , ($r = -0.761$, $p = 0.000$) (**Figure 3c**). There was a significant
411 positive correlation between %mud and sediment C_{stocks} ($r = 0.761$, $p = 0.004$) (**Figure 3d**). The
412 association between pore water pH levels and sediment C_{stock} represented a significant negative
413 relationship, ($r = -0.545$, $p = 0.003$) (**Figure 3e**). There was a significant positive relationship
414 between salinity and sediment C_{stocks} ($r = 0.876$, $p = 0.000$) (**Figure 3f**) and between elevation and
415 sediment C_{stock} ($r = 0.719$, $p = 0.000$) (**Figure 3g**). The concentration of sulphates was significantly

Fig. 3 Relationships between sediment C_{stocks} and (a) dry bulk density (DBD) including linear regression line and equation, R^2 and n values. Spearman's rho relationships between sediment C_{stocks} and (b) mean grain size, (d) %mud and (g) sulphate levels represented by $\text{Log}_{10}(X)$ values, including r, p and n values. Pearson's relationship between sediment C_{stocks} and (c) degree of sorting, (e) pH, (f) salinity, and (h) elevation, (i) above-ground biomass, including r, p and n values, for all sites: Creek Rythe, Cowes, Farlington Marshes, Hayling Island, Porchester and Ryde

416 negatively correlated with sediment C_{stock} ($r = -0.522$, $p = 0.004$) (**Figure 3h**), but concentration of
417 nitrites was only moderately significantly correlated to sediment C_{stock} ($r = -0.423$, $p = 0.031$).
418 There was no significant relationship between sediment C_{stock} and leaf density ($r = -0.095$, $p =$
419 0.616), but above-ground biomass was significantly related to sediment C_{stocks} ($r = 0.595$, $p =$
420 0.001) (**Figure 3i**).

421 The multivariate relationship between sediment C_{stock} and the predictor variables was
422 explained in a PLS regression model. The cumulative fraction (R-sq) of the first four components
423 (DBD, Sorting, %Mud and pH) was 0.90, showing a high degree of determination, meaning they
424 explain 90% of the variation in the model data. According to the model, the most important factors
425 responsible for the variation in sediment C_{stocks} were DBD, followed in rank order by sorting
426 coefficient, %mud, pH, sulphates, nitrites, elevation, salinity, above-ground biomass, mean grain
427 size and leaf density (**Figure 4**). DBD, sorting coefficient, %mud, sulphates, elevation, mean grain
428 size, and leaf density showed negative relationships with sediment C_{stocks} , while pH, nitrites, above-
429 ground biomass and salinity were positively related to sediment C_{stocks} .

430 **Fig. 4** Partial least square regression model coefficient plot. The model assesses the relative influence of different predictors in
431 sediment C_{stocks} . Predictors are dry bulk density (DBD), sorting coefficient, %mud, pH, nitrites (NO_2^-), sulphates (SO_4^{2-}), elevation,
432 salinity, above-ground biomass, mean grain size, and leaf density, ranked by level of importance from left, most important, to right,
433 least important

432 The six predictors indicated by PLS as main drivers for variation in sediment C_{stocks} , with a
433 coefficient > 5 , were selected for inclusion in the PCA, namely DBD, sorting coefficient, %mud,
434 pH, sulphates and nitrites. The first two principal components cumulatively explained 78.3% of
435 the variation in the data, with eigenvalues > 1 . The first principal component (PC1) explained
436 62.6% of the variation, with 15.7% being explained by the second principal component (PC2).
437 Sediment characteristics showed the largest influence on PC1, with DBD and sorting coefficient

438 with positive loadings on the component, closely related to sediment C_{stock} and %mud with negative
439 loadings (**Figure 5**). This indicates that PC1 primarily expresses sediment characteristics. PC2 was
440 mainly negatively influenced by the concentration of nitrites (**Figure 5**). The relationship between
441 sediment C_{stocks} and the different variables was not uniform between sampling sites, with points
442 from Ryde, Cowes and Farlington Marshes grouped separately from other sites, while points from
443 Creek Rythe, Hayling Island and Porchester were closer together (**Figure 5**). Points from Creek
444 Rythe and Hayling Island had low values for both PC1 and PC2, representing sites with highest
445 C_{stock} , high %mud and lowest DBD. Points for Ryde had high positive values for PC1 and three
446 points had high negative values for PC2, with the other 2 points exhibiting low values for PC2. Ryde
447 was the site with lowest C_{stock} , lowest %mud, highest DBD, highest degree of sorting and higher
448 concentration of both pore water nutrients (nitrites and sulphates). Cowes and Farlington Marshes
449 had positive relationships with both principal components, representing sites with similar sorting,
450 %mud and pH levels. However, Cowes had higher values for PC1, while points from Farlington
451 Marshes was grouped closer to the centre of both axis. This could be explained by the higher values
452 for C_{stock} and lower DBD in Farlington Marshes (**Figure 5**).

453

Fig. 5 Principal component analysis (PCA) showing the six seagrass study sites, Creek Rythe, Cowes, Farlington Marshes, Hayling Island, Porchester and Ryde, related to the five most relevant predictor variables, dry bulk density (DBD), sorting coefficient, %mud, pH, nitrites (NO_2^-) and sulphates (SO_4^{2-}) (Figure 4) in terms of sediment C_{stocks} as the response variable

454

455 **Discussion**

456 Results from this study provide valuable insights into the environmental factors influencing
457 sediment carbon storage in temperate seagrass ecosystems, specifically within intertidal seagrass
458 meadows in the Solent region, England. Seagrass meadows with different species composition,

459 showed significantly different sediment C_{stocks} , dry bulk density, mean grain size, degree of sorting,
460 proportion of mud and pore water pH and salinity. Overall, the main factors significantly related
461 to seagrass sediment C_{stocks} were: elevation, pore water sulphates, pH and salinity; and sediment
462 degree of sorting, grain size, proportion of mud and dry bulk density. Indeed, dry bulk density can
463 be used as a predictor for sediment C_{stocks} as indicated by the regression equation developed by this
464 study. However, when the combined association of factors was analysed, dry bulk density, sorting
465 coefficient, proportion of mud, pH, nitrites and sulphates showed the greatest influence on sediment
466 C_{stocks} . Moreover, sediment characteristics such as dry bulk density, sorting coefficient and
467 proportion of mud, acted in conjunction to explain the bulk of the variation in sediment C_{stocks} .

468 There was no significant difference in leaf density between sampling sites.
469 However, meadows with mixed species, and dense continuous beds, namely Creek Rythe, Hayling
470 Island and Porchester, supported higher sediment C_{stocks} than sites that formed beds with single
471 species, such as Farlington Marshes and Ryde, or were very patchy as at Cowes. Higher species
472 diversity, can increase seagrass efficiency in reducing currents and consequent sediment
473 resuspension, therefore contributing to organic matter deposition, especially in species such as
474 *Zostera spp.* and *Thalassia spp.*, with blade-like leaves (Verduin and Backhaus, 2000; Koch *et al.*,
475 2006; Peralta *et al.*, 2008; Hendricks *et al.*, 2008, Mazarrasa *et al.*, 2018). Additionally, dense and
476 continuous meadows retain more autochthonous carbon, such as leaf detritus, which combined with
477 their ability to accumulate finer sediment particulates, enhances their carbon storage and
478 sequestration capacity (Miyajima *et al.*, 2017; Oreska *et al.*, 2017; Rocart *et al.*, 2017; Mazarrasa
479 *et al.*, 2018).

480 Recent studies assessing the role of environmental parameters in determining seagrasses'
481 carbon sink potential corroborate the results of this research, with sediment properties being

482 identified as highly influential (Lavery *et al.*, 2013; Duarte *et al.*, 2011; Dahl *et al.*, 2016; Mazarrasa
483 *et al.*, 2018; Rohr *et al.*, 2018). Results from this study show that large sediment C_{stocks} are strongly
484 linked to a high proportion of mud and low bulk density, suggesting that seagrass meadows with
485 such sediment characteristics have a higher potential as natural carbon sinks (Dahl *et al.*, 2016;
486 Mazarrasa *et al.*, 2018; Rohr *et al.*, 2018). Grain size is strongly related to sediment porosity and
487 density, which are important factors regulating oxygen concentrations in the sediment and
488 consequent degradation of organic matter by microbial activity (Benner *et al.*, 1984; Enriquez *et*
489 *al.*, 1993; Deming and Harass, 1993; Dahl *et al.*, 2016). Seagrass meadows, especially ones with a
490 low contribution of autochthonous carbon sources to sediment pools, can increase the concentration
491 of fine grain particles in the sediment by reducing water velocity and facilitating sedimentation,
492 thus promoting high carbon storage (Serrano *et al.*, 2016). Higher proportions of fine grains in mud
493 substrate, with greater particle surface areas, also contribute to the preservation and accumulation
494 of organic matter (Mayer, 1994; Dahl *et al.*, 2016; Mazarrasa *et al.*, 2018; Rohr *et al.*, 2018). Sites
495 with smaller grain size fractions, such as Hayling Island and Creek Rythe, are likely to have lower
496 soil permeability and more anoxic conditions as a result of smaller interstitial spaces, which reduces
497 organic matter degradation rates by decreasing oxygen exchange and redox potential, contributing
498 to their higher sediment C_{stocks} (Hedges, 1995; Wilson *et al.*, 2008; Dahl *et al.*, 2016). The
499 relationship between carbon storage and sediment characteristics is more evident in meadows with
500 relatively low seagrass biomass and a high proportion of finer particle sizes, such as the temperate
501 ones studied here, while in meadows dominated by seagrass species with a greater biomass, e.g.
502 *Posidonia spp.*, the amount of autochthonous carbon seems to be more important for carbon storage
503 than mud and silt content (Serrano *et al.*, 2016).

504 Dahl *et al.* (2016) demonstrated that sediment density has a negative relationship with
505 sediment organic carbon, affirming the results for the variance in dry bulk density among sites in

506 this study, where sites with higher bulk density (and therefore lower porosity) had lower sediment
507 C_{stocks} , e.g. Ryde and Cowes. In combination with particle size, the degree of exposure to
508 hydrodynamic forces such as waves, tides and currents, is also a determinant factor for
509 sedimentation and erosion in coastal areas (Maxwell *et al.*, 2017; Mazarrasa *et al.*, 2018). The level
510 of exposure in seagrass meadows is usually reflected in the proportion of fine sediment particles
511 (e.g. % mud), being higher in sheltered areas compared to more exposed sites (Van Keulen and
512 Borowitzka, 2003; Mazarassa *et al.*, 2018). This is likely to be the case for meadows at Ryde and
513 Cowes on the Isle of Wight, which are more exposed to wave activity, and present lower sediment
514 C_{stocks} , than the mainland sites. The degree of sorting can also be used as a proxy to indicate physical
515 exposure related to movement of water masses, with better sorted particles representing slower
516 deposition driven by stable hydrodynamic conditions (Folk and Ward, 1957; Mazarrasa *et al.*,
517 2017; Rohr *et al.*, 2018). All sites in this study had sediments classified as very well sorted, but the
518 two on the Isle of Wight had significantly higher degrees of sorting and lower sediment C_{stock} values
519 than the other sites. These results conform with those obtained by Rohr *et al.* (2018) for seagrasses
520 in the Baltic, where exposure to wave activity was an important driver for sediment C_{stock} in *Z.*
521 *marina* meadows, with high exposure leading to lower sediment C_{stock} due to the potential export
522 of carbon to other adjacent ecosystems. Other factors related to sedimentation processes have
523 previously also shown to be relevant to seagrass sediment carbon storage, such as water depth, in-
524 situ primary productivity, sedimentation rates and the trapping of fine-grained sediment and
525 organic matter (Serrano *et al.*, 2016; Dahl *et al.*, 2016; Mazarrasa *et al.*, 2018).

526 Even though there was a limited range of pore water pH values across all six sites, they
527 were significantly higher on sites with less sediment C_{stocks} , and showed a significant negative
528 relationship with seagrass sediment carbon storage. This association is supported by Ivers *et al.*'s.,
529 (1997) findings of a decrease in photosynthetic rates when pH increased by 0.6 units for *Posidonia*

530 *oceanica* and *Cymodocea nodosa*, and 0.8 units for *Z. noltei*. Furthermore, Egea *et al.* (2018) found
531 no effect on seagrass production with increased acidification, reporting a slight increase in carbon
532 stocks with lower pH levels. At a broader scale, studies suggest an increase in seagrass productivity
533 and consequent carbon storage in acidic scenarios, with Garrand and Beaumont (2014) quantifying
534 that the reduction in pH of ocean surface waters is expected to enhance both above- and
535 belowground biomass, leading to an 82–94% increase in seagrass carbon storage and sequestration
536 potential, potentially increasing ocean storage of carbon by 12–14%.

537 It has been suggested that the amount of organic carbon in seagrass sediments can be positively
538 linked to above and belowground plant productivity, with seagrass productivity being sensitive to
539 nutrient input, often decreasing substantially as a result of light limitation during algal blooms when
540 high nutrient loadings occur (Hauxwell *et al.*, 2001; Schmidt *et al.*, 2012; Burkholder *et al.*, 2007;
541 Kirwan and Mudd, 2012; Armitage and Fourqurean, 2016). Nutrient availability in seagrass
542 sediments is closely related to microbial activity in their rhizosphere, associated with the release
543 of oxygen into the sediment by seagrasses' rhizomes, enhancing bacterial activity and nitrogen
544 fixation at depths that would otherwise be anoxic (Perry and Dennison, 2000). Thus, sediment grain
545 size might be one of the limiting factors in nutrient cycling, with fine grain sediments having lower
546 concentrations of oxygen with depth, decreasing microbial activity (Mazarrasa *et al.*, 2018). An
547 increase in microbial activity can potentially explain the high concentration of pore water nitrites
548 and sulphates in the upper layer of sites in this study with lower sediment C_{stocks} , such as at Ryde
549 and Cowes.

550 Elevation in relation to mean sea level was correlated with seagrass sediment C_{stocks} . Ryde and
551 Cowes, which were located at the lowest elevations in relation to mean sea level, had the lowest
552 values of sediments C_{stocks} . Although all sites in this study were intertidal, differences in elevation
553 are related to varying periods of emersion and desiccation between low and high tide, which

554 impacts carbon sequestration processes, since areas with higher exposed periods could have higher
555 rates of photosynthesis, therefore sequestering and storing more carbon in their sediments (Short
556 and Neckles, 1999; Mazarrasa *et al.*, 2018). Conversely, some studies have identified increased
557 desiccation stress as a factor that slows recovery time in intertidal seagrass beds growing at higher
558 elevations, indicating higher vulnerability to extreme weather events linked to climatic change (de
559 Fouw *et al.*, 2016; El Hacen *et al.*, 2018).

560 In conclusion, this study showed that seagrass meadows within the same climatic region, and
561 locality, do not share the same potential for long-term sediment carbon storage, and that
562 environmental characteristics strongly influence this ecosystem service. Therefore, while seagrass
563 research at global, and even continental, scales are extremely important, caution must be taken with
564 extrapolations and generalisations across different regions. These results also show that individual
565 seagrass meadows might not be representative of the whole ecosystem, and highlights the need for
566 the assessment and consideration of multiple environmental features and their interactions in
567 seagrass blue carbon research, such as dry bulk density, sorting coefficient and other sediment
568 characteristics, above-ground biomass, nutrients, elevation in relation to mean sea level etc.. In the
569 Solent estuary, larger stocks were associated with meadows located in sheltered bays, with high
570 sediment mud content and well sorted particles. Conversely, exposed meadows subject to intense
571 anthropogenic disturbance are likely to experience a decline in their capacity to sequester and store
572 carbon in the long-term, as shown by sites with patchy seagrass landscapes, high surface nutrient
573 levels and lower sediment C_{stocks} . This indicates that the most influential factors driving temperate
574 seagrass sediment C_{stocks} , namely dry bulk density, degree of sorting, and proportion of mud, should
575 be monitored in conjunction with pore water sulphates, pH and salinity, elevation and mean grain
576 size, in conservation and restoration projects that aim to promote the carbon sink potential of
577 intertidal seagrass ecosystems. It is also evident that seagrass carbon sink potential is regulated by

578 a combination of multiple environmental factors, encompassing sediment and vegetation variables,
579 highlighting the potential vulnerability of these ecosystems to climate change, such as sea level
580 rise. Therefore, key factors should be considered, individually or ideally in combination, when
581 developing and implementing conservation or restoration projects, and climate change mitigation
582 strategies, using seagrass ecosystems.

Table 2

Study sites	Coordinates	Meadow extent (ha)*	C _{Stock} (30cm)** (MgCha ⁻¹)	Vegetation	Leaf density (m ⁻²)	Above-ground Biomass (MgC ha ⁻¹)	Elevation (m)	Dry bulk density (gdm ⁻³)	Mean grain size (µm)	Sorting coefficient (φ)	% Mud	pH	Salinity (‰)	NO ₂ ⁻ (µmolL ⁻¹)	SO ₄ ⁻² (µmolL ⁻¹)
Hayling Island	50°47'54"N, 0°59'48"W	100.24	51.13 ± 7.84 (A)	<i>Z. marina</i> / <i>Z. angustifolia</i> / <i>Z. noltei</i> / <i>Ruppia</i> spp. Dense	336.7 ± 95.0	0.37 ± 0.13 (AB)	-0.81 ± 0.13 (C)	0.72 ± 0.04 (CD)	20.37 ± 3.28 (C)	-1.73 ± 0.07 (BC)	93.8 ± 7.07 ± 3.2 (A)	15.6 ± 0.9 (A)	18914 ± 9765	488.9 ± 199.2	
Creek Rythe	50°49'3"N, 0°53'33"W	70.1	45.31 ± 3.53 (A)	<i>Z. marina</i> / <i>Z. angustifolia</i> / <i>Z. noltei</i> / <i>Ruppia</i> spp. Dense	367.0 ± 115.1	0.50 ± 0.25 (A)	0.0008 ± 0.06 (A)	0.68 ± 0.08 (D)	24.25 ± 10.82 (C)	-1.78 ± 0.09 (CD)	87.5 ± 7.16 ± 3.6 (A)	12.8 ± 2.8 (A)	18538 ± 15929	554 ± 254	
Porchester	50°50'13"N, 1°7'51"W	94.92	45.23 ± 12.10 (A)	<i>Z. angustifolia</i> / <i>Z. noltei</i> Patchy	302.0 ± 76.1	0.32 ± 0.07 (ABC)	-0.55 ± 0.11 (B)	0.89 ± 0.20 (BC)	55.23 ± 31.09 (BC)	-1.99 ± 0.19 (D)	79.3 ± 7.17 ± 13.0 (AB)	12.6 ± 1.9 (A)	20363 ± 19780	424.8 ± 154.6	
Farlington Marshes	50°50'2"N, 1°2'24"W	31.2	40.23 ± 4.16 (A)	<i>Z. angustifolia</i> Very patchy	584 ± 427	0.25 ± 0.14 (ABC)	0.059 ± 0.06 (A)	0.94 ± 0.09 (B)	48.71 ± 29.08 (BC)	-1.76 ± 0.18 (BCD)	79.7 ± 7.59 ± 6.3 (AB)	14.4 ± 3.6 (A)	10383 ± 9109	2334 ± 2317	
Cowes	50°45'55"N, 1°16'56"W	27.1	14.22 ± 7.59 (B)	<i>Z. marina</i> / <i>Z. noltei</i> Very patchy	346 ± 247	0.18 ± 0.07 (BC)	-1.82 ± 0.05 (E)	1.27 ± 0.02 (A)	72.40 ± 36.91 (B)	-1.51 ± 0.16 (B)	66.62 ± 18.5 (B)	7.42 ± 0.34 (AB)	2.0 ± 0.0 (B)	6271 ± 5892	1011 ± 985
Ryde	50°44'02"N, 1°09'23"W	82.47	6.65 ± 1.73 (B)	<i>Z. angustifolia</i> Patchy	427 ± 430	0.07 ± 0.03 (C)	-1.48 ± 0.11 (D)	1.46 ± 0.03 (A)	225.01 ± 7.78 (A)	-0.48 ± 0.02 (A)	1.14 ± 7.66 ± 0.03 (C)	1.4 ± 0.6 (B)	69509 ± 69739	4755 ± 5628	

584 *Area derived from Marsden and Chesworth (2015).

585 ** Cores from Cowes (CWST) were 20cm deep.

586

587

588

589 **Supplementary Table**

590 **Table A:** Summary of ANOVA results for all environmental variables between the six sampling
 591 sites. Mathematical transformation ($\text{Log}_{10}(X)$) performed when assumptions of normality of
 592 residuals were not met – AD (p) is the Anderson- Darling normality test result ($p < 0.05$).

Variables	df	F	p	R-sq	Pooled StD	AD (p)	Transformation
Sediment C_{stock} (Mg C Ha ⁻¹)	29	34.70	0.000	0.8785	7.05	0.438 (0.277)	Normal
Dry bulk density (g dm ⁻³)	29	51.08	0.000	0.9141	0.10	1.136 (<0.05)	$\text{Log}_{10}(X)$
Mean grain size (μm)	29	67.04	0.000	0.9229	21.65	1.242 (<0.05)	$\text{Log}_{10}(X)$
Sorting coefficient (ϕ)	29	92.62	0.000	0.9411	0.13	0.613 (0.103)	Normal
% Mud	29	17.56	0.000	0.7853	9.88	1.348 (<0.05)	$\text{Log}_{10}(X)$
pH	29	5.49	0.002	0.5664	0.28	1.090 (0.006)	Normal
Salinity (‰)	29	33.58	0.000	0.8888	2.20	0.628 (0.091)	Normal
leaf density (leaf/m ⁻²)	29	0.68	0.642	0.1243	68.91	0.896 (0.019)	Normal
Above-ground Biomass (Mg C Ha ⁻¹)	29	5.97	0.001	0.5543	0.14	0.290 (0.587)	Normal
Nitrites NO_2^- (μmolL^-)	29	1.36	0.280	0.2542	0.57	1.757 (<0.05)	$\text{Log}_{10}(X)$
Sulphates SO_4^{2-} (μmolL^-)	29	3.29	0.026	0.4641	0.39	2.717 (<0.05)	$\text{Log}_{10}(X)$
Elevation (m)	29	332.09	0.00	0.9858	0.09	0.142 (0.968)	Normal

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